

chapter overview

- Ocean currents are water masses in motion
- Surface currents are influenced by major wind belts
- Currents redistribute global heat
- Locally, surface currents affect coastal climates
- Thermohaline circulation affects deep currents
- Currents influence marine life by affecting the growth of microscopic algae (basis of most oceanic food webs)

Types of Ocean Currents

- Surface currents
 - Wind-driven
 - Primarily horizontal motion
- Deep currents
 - Driven by differences in density caused by differences in temperature and salinity
 - Vertical and horizontal motions

7.1 – How Are Ocean Currents Measured?

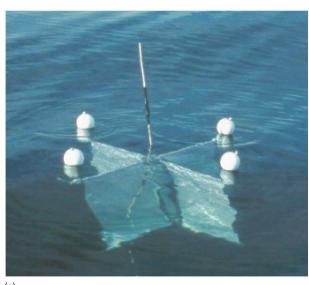
Measuring Surface Currents

Direct methods

- Floating devices whose position is tracked through time
- Fixed current meter (such as a propeller flow meter)

Indirect methods

- Pressure gradients (water flows downhill)
- Radar altimeters (dynamic topography maps)
- Doppler flow meter (use of low-frequency sound signals)

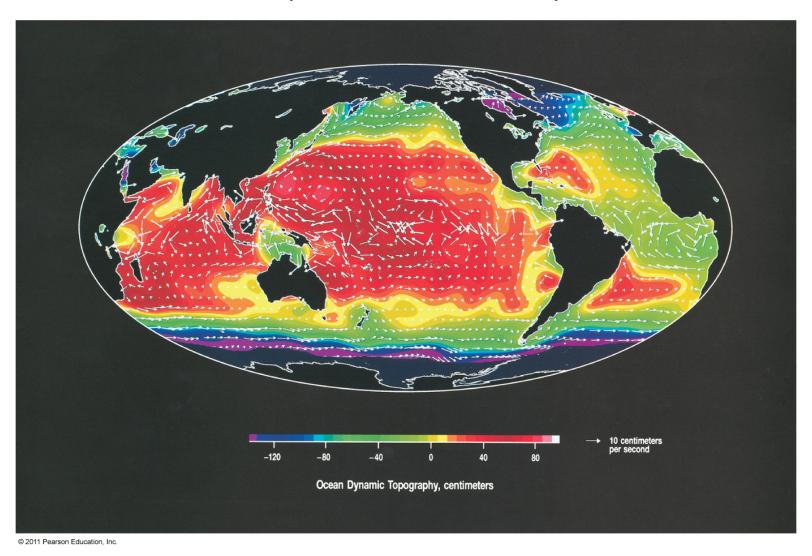






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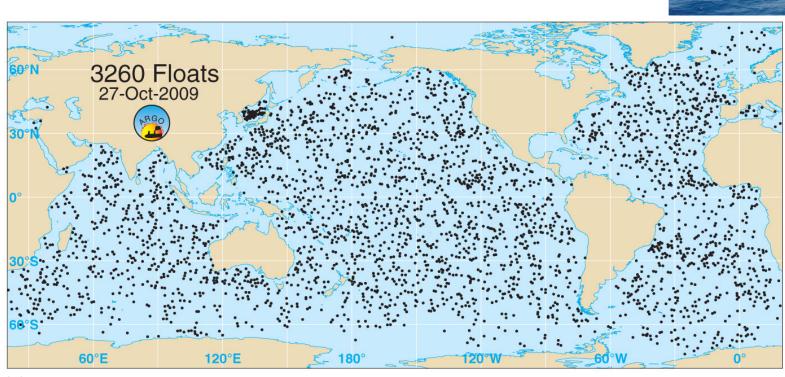
Satellite View (from radar altimeter data) of Global Surface Current Flow, September 1992 to September 1993



Measuring Deep Currents

- It is very difficult to track deep currents
 - Argo program
 - release of thousands floating device that sink to depth, record data, come back to surface, transmit, sink again
 - Identification of distinctive temperature and salinity of deep water masses
 - Chemical tracers
 - Natural
 - Intentionally added
 - Unintentionally added

Argo program

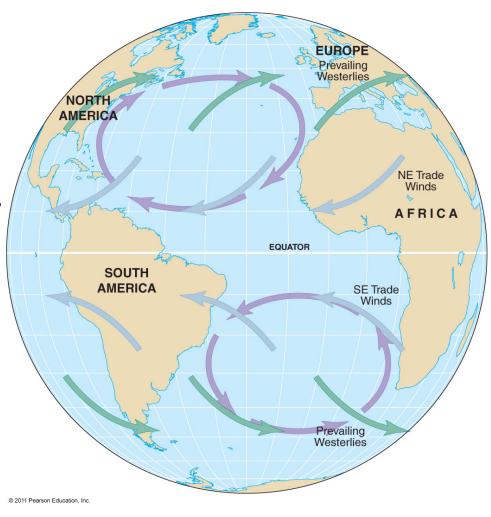


7.2 – What Creates Ocean Surface Currents and How Are They Organized?

- Occur within and above pycnocline, because of frictional drag between wind and ocean
 - to a depth of about 1 km
 - affect only about 10% of world's ocean water
- Affected mostly by major wind belts of the world, so they generally follow wind belt patterns
- Other factors:
 - Distribution of continents and geometry of ocean basins
 - Gravity
 - Seasonal changes
 - Coriolis effect
 - Friction

general surface currents patterns

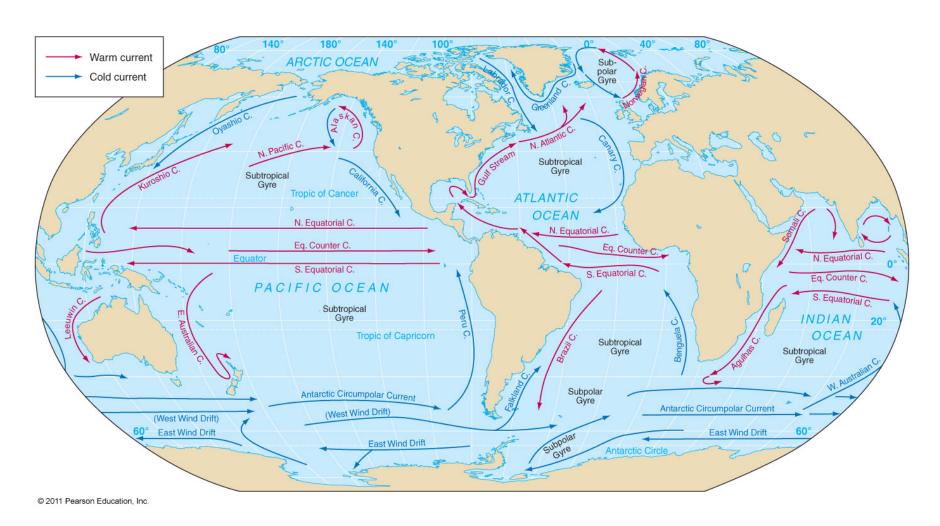
- Only 2% of wind energy is transmitted to ocean
- On a planet without land masses, ocean currents would follow wind belts, creating large, circular loops (gyres) of moving water
- On a planet with continents, currents are deflected by land
- Example from the Atlantic Ocean



Subtropical Gyres

- There are five Subtropical Gyres
 - North and South Atlantic, North and South Pacific, Indian
 - rotate clockwise in Northern Hemisphere, counterclockwise in the Southern Hemisphere
 - average drift time is between 3 and 6 years
- Gyres are bounded by:
 - Equatorial current
 - Western Boundary currents
 - Northern or Southern Boundary currents
 - Eastern Boundary currents
- Centered around 30° latitude

Subtropical Gyres and Currents



the Four Currents that make up a Subtropical Gyre

All gyres have four main currents flowing into one another:

- Equatorial Currents
 - North or south of the Equator, traveling westward along it
- Western Boundary Currents warm waters
- Northern or Southern Boundary Currents easterly water flow across ocean basin
- Eastern Boundary Currents cool waters

Gyres and Boundary Currents

TABLE 7.1

SUBTROPICAL GYRES AND SURFACE CURRENTS

^bDenotes a western boundary current of a gyre, which is relatively *fast*, *narrow*, and *deep* (and is also a *warm-water* current).

| | North Pacific (Turtle) Gyre | | North Atlantic (Columbus) Gyre | | Indian Ocean (Majid) Gyre |
|---------------|---------------------------------------|----------------|---------------------------------|--------|--------------------------------------|
| | North Pacific Current | | North Atlantic Current | | South Equatorial Current |
| | California Current ^a | | Canary Current ^a | | Agulhas Current ^b |
| | North Equatorial Current | | North Equatorial Current | | West Wind Drift |
| | Kuroshio (Japan) Current ^b | | Gulf Stream ^b | | West Australian Current ^a |
| | South Pacific (Heyerdahl) Gyre | | South Atlantic (Navigator) Gyre | | Other Major Currents |
| an | South Equatorial Current | ean | South Equatorial Current | an | Equatorial Countercurrent |
| 000 | East Australian Current ^b | 000 | Brazil Current ^b | Ocean | North Equatorial Current |
| Pacific Ocean | West Wind Drift | Atlantic Ocean | West Wind Drift | Indian | Leeuwin Current |
| Pa | Peru (Humboldt) Current ^a | Atl | Benguela Current ^a | In | Somali Current |
| | Other Major Currents | | Other Major Currents | | |
| | Equatorial Countercurrent | | Equatorial Countercurrent | | |
| | Alaskan Current | | Florida Current | | |
| | Oyashio Current | | East Greenland Current | | |
| | | | Labrador Current | | |
| | | | Falkland Current | | |

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Other Surface Currents

Equatorial Countercurrents

- equatorial currents push water westward
- an eastward flow (between North and South Equatorial Currents) develops because of a higher sea level to the west
- different in the three oceans

Subpolar Gyres

- Rotate opposite subtropical gyres
- Smaller and fewer than subtropical gyres

Other factors affecting ocean surface circulation

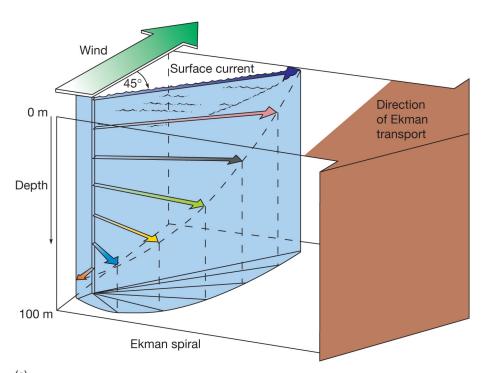
- Ekman Spiral
- Ekman transport

Geostrophic currents

Western intensification of subtropical gyres

Ekman Spiral

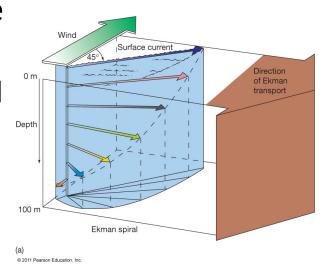
- Surface currents move at an angle to the wind
- To the right in the NH, to the left in the SH
- Upper layer drags lower layer, but with less energy
- Each successive layer moves increasingly to the right in the Northern Hemisphere
 - Coriolis effect
- Movement stops at about 100 m of depth
- The Ekman spiral describes speed and direction of seawater flow at different depths.

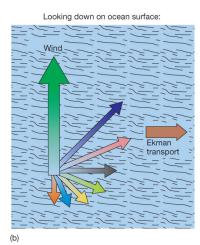


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Ekman Transport

- Average movement of seawater under influence of wind
- 90° to right of wind in NH
- 90° to left of wind in SH
- "Ideal conditions" rarely exist, so on average Ekman transport is 70°
- In coastal waters, might be 0°

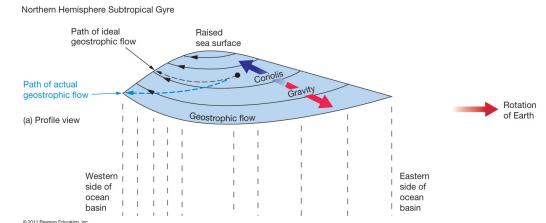


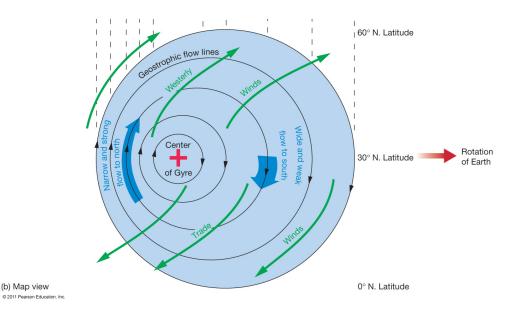


Geostrophic Currents

(b) Map view

- Ekman transport piles up water within subtropical gyres.
- Surface water flows downhill and to the right.
- Geostrophic flow: balance of Coriolis Effect and gravitational forces
- Ideal geostrophic flow
- Friction generates actual geostrophic flow





Western Intensification

- The top of the hill of water is displaced toward the west due to Earth's rotation
- Western boundary currents are intensified in both hemispheres
 - Faster, Narrower, Deeper, Warm
- Coriolis Effect contributes to western intensification
- Eastern boundary currents tend to have the opposite properties of Western Boundary Currents
 - Cold, Slow, Shallow, Wide

Eastern and Western Boundary Currents

| TABLE 7.2 | CHARACTERISTICS OF WESTERN AND EASTERN BOUNDARY CURRENTS OF SUBTROPICAL GYRES |
|-----------|---|
| | |

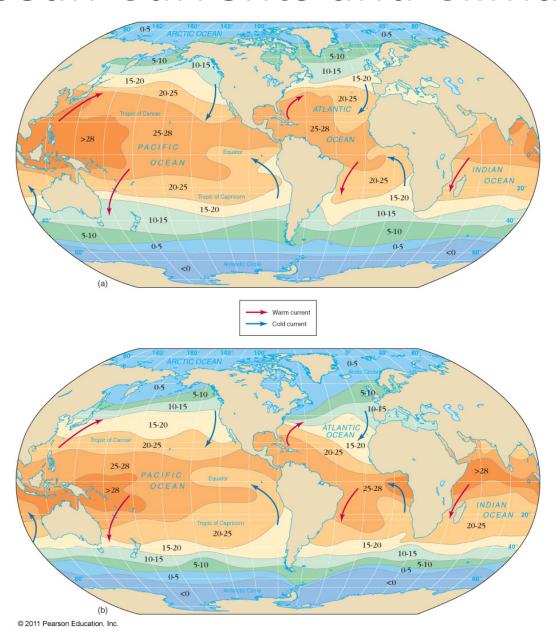
| Current type | Examples | Width | Depth | Speed | Transport volume (millions of cubic meters per second ^a) | Comments |
|--------------------------------|--|---|--|--|--|---|
| Western boundary current | Gulf Stream, Brazil Current, Kuroshio Current | Narrow: usually less than 100 kilometers (60 miles) | Deep: to depths of 2 kilometers (1.2 miles) | Fast: hundreds of kilometers per day | Large: as much as 100 Sv ^a | Waters derived from low latitudes and are warm; little or no upwelling |
| Eastern boundary current | Canary Current, Benguela Current, California Current | Wide: up to 1000 kilometers (600 miles) | Shallow: to depths of 0.5 kilometer (0.3 mile) | Slow: tens of kilometers per day | Small: typically 10 to 15 Sv ^a | Waters derived from middle lati- tudes and are cool; coastal upwelling common |

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Ocean Currents and Climate

- Warm coastal ocean currents warm the air above. Air absorbs more moisture
 - Warm, humid air
 - Humid climate on adjoining landmass
 - Florida
- Cool coastal ocean currents cool the air above. Air cannot acquire moisture
 - Cool, dry air
 - Dry climate on adjoining landmass
 - California

Ocean Currents and Climate



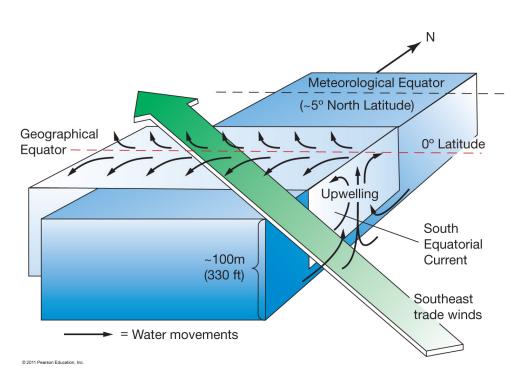
7.3 – What Causes Upwelling and Downwelling?

- Upwelling Vertical movement of cold, nutrientrich water to surface
 - High biological productivity
- Downwelling Vertical movement of surface water downward in water column
 - Low productivity, but downwelling carries oxygen to deep waters
- Upwelling and downwelling provide important mixing mechanisms between surface and deep waters

How do upwelling and downwelling occur?

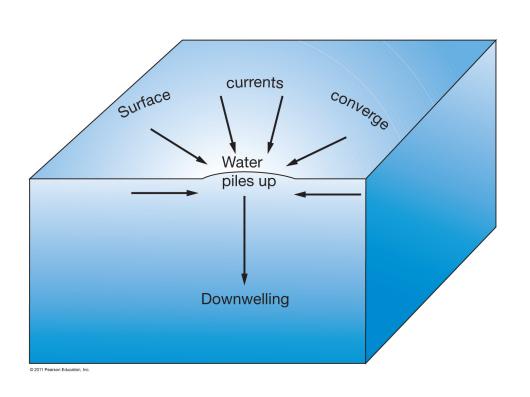
- Diverging surface water
- Converging surface water
- Coastal upwelling and downwelling
- Offshore winds
- Seafloor obstructions
- Sharp bends in coastline
- Absence of pycnocline in high-latitude regions

Diverging Surface Water



- Surface waters move away from an area on the ocean's surface
- Example: Trade Winds along the equator
 - Ekman transport to the right in NH
 - Ekman transport to the left in SH
- Equatorial upwelling

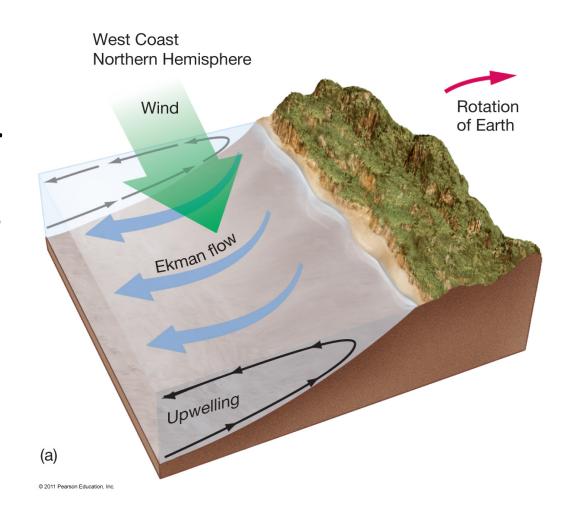
Converging Surface Water



- Occurs when surface waters move toward each other
- Water piles up and has no place to go but downward (downwelling)
- Low biological productivity

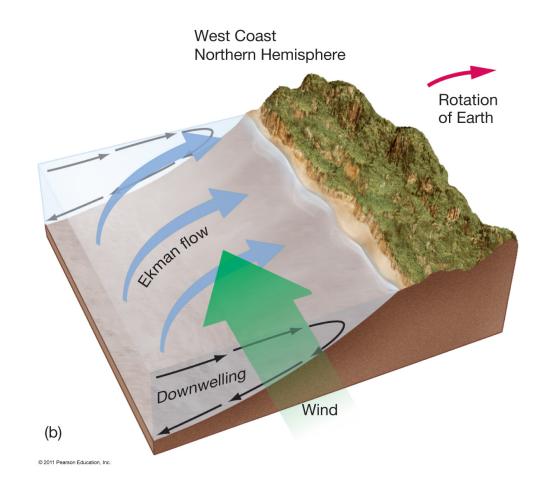
Coastal Upwelling

- Ekman transport moves surface seawater offshore.
- Cool, nutrient-rich deep water comes up to replace displaced surface waters.
- Example: U.S.
 West Coast



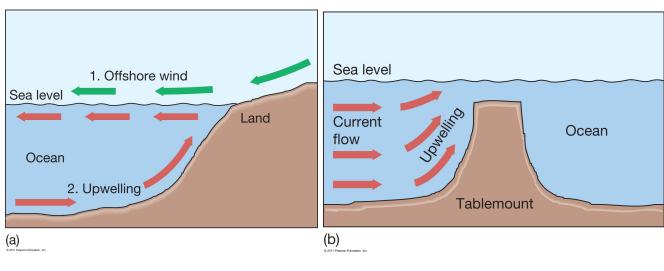
Coastal Downwelling

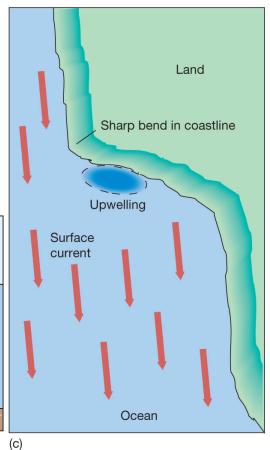
- Ekman transport moves surface seawater toward shore.
- Water piles up, has no place to go, and moves downward in water column
- Lack of marine life



Other Causes of Upwelling

- Offshore winds
- Seafloor obstruction
- Coastal geometry change
- Absence of a pycnocline





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7.4 – What Are the Main Surface Circulation Patterns in Each Ocean Basin?

- The specific pattern of surface currents varies from ocean to ocean, depending on:
 - the geometry of the ocean basin
 - the pattern of major wind belts
 - seasonal factors, and other periodic changes
- We are going to look at the surface ocean water circulation of:
 - the Antarctic Ocean
 - the Atlantic Ocean
 - The Indian Ocean
 - the Pacific Ocean

Antarctic Circulation

 Antarctic circulation is dominated by movement of water masses in the southern Atlantic, Indian, and Pacific Oceans, at latitudes south of 50°S

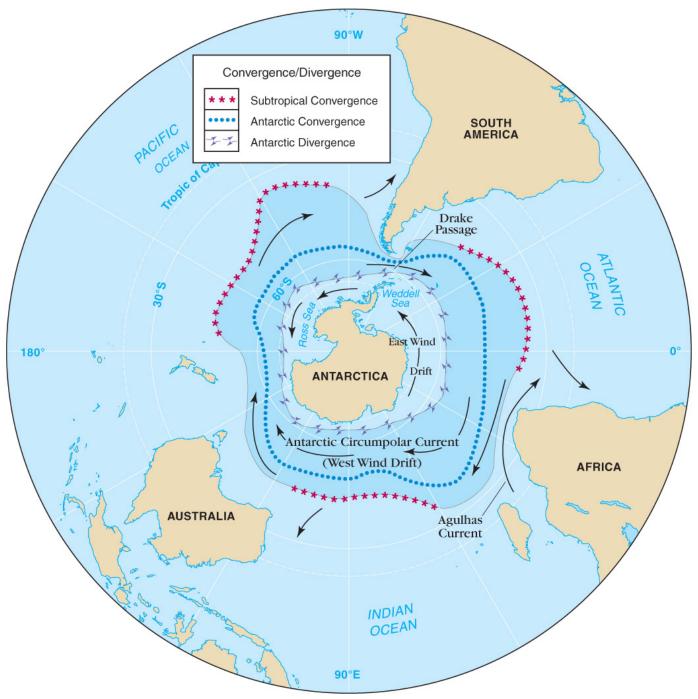
- There are two main currents:
 - Antarctic Circumpolar Current (West Wind Drift)
 - East Wind Drift

Antarctic Circumpolar Current (West Wind Drift)

- Only current to completely encircle Earth
- Powered by the prevailing westerly wind belt
- Moves more water than any other current
- Subtropical convergence located at 40°S

East Wind Drift

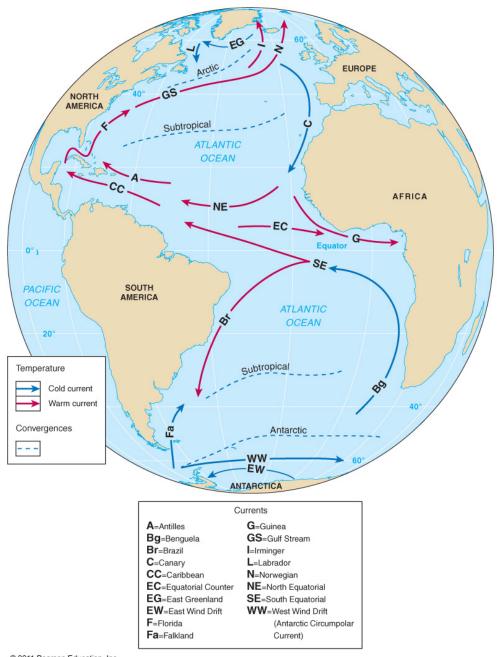
- propelled by Polar Easterlies
- Creates surface divergence with opposite flowing West
 Wind Drift



- These two currents create zones of convergence and divergence
- Antarctic Convergence (or Antarctic Polar Front)
 - Cold, dense Antarctic waters converge with warmer, less dense sub-Antarctic waters, at about 50°S
 - Downwelling
 - Northernmost boundary of Antarctic Ocean
- Antarctic Divergence
 - Upwelling
 - Nutrients from the bottom, and as a consequence,
 Abundant marine life

Atlantic Ocean Circulation

- Consists of two large subtropical gyres:
 - North AtlanticSubtropical Gyre
 - South AtlanticSubtropical Gyre



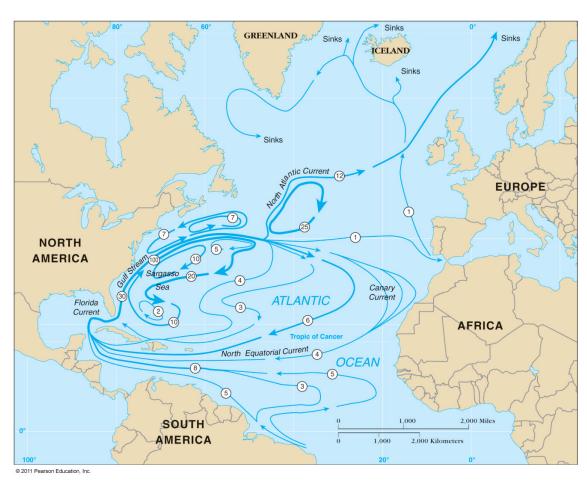
North and South Atlantic Subtropical Gyres

- North Atlantic
 Subtropical Gyre
 - North Equatorial Current
 - Gulf Stream
 - North Atlantic Current
 - Canary Current
 - South Equatorial Current
 - Atlantic Equatorial
 Counter Current

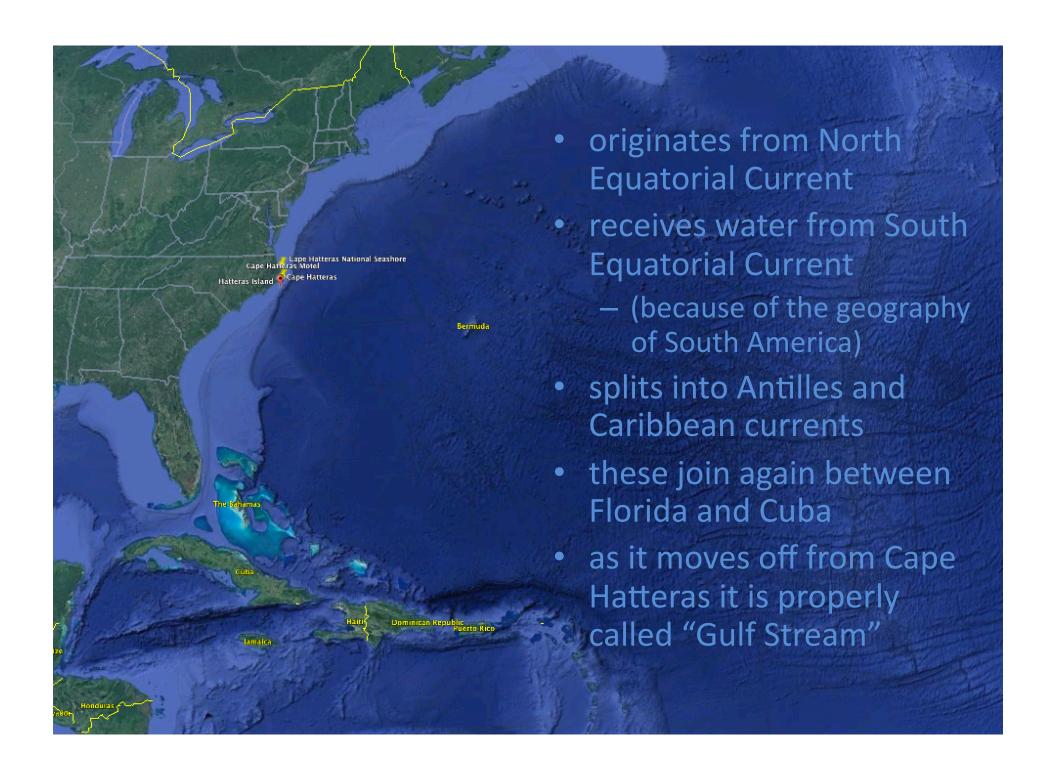
- South Atlantic
 Subtropical Gyre
 - Brazil Current
 - Antarctic Circumpolar
 Current
 - Benguela Current
 - South Equatorial Current
 - Falkland Current

Gulf Stream

- moves
 northward along
 the east coast of
 North America
- warms coastal states
- warms northern Europe



Numbers indicate average flow rates in Sverdrups One Sverdrup is equivalent to 1,000,000 m³/sec

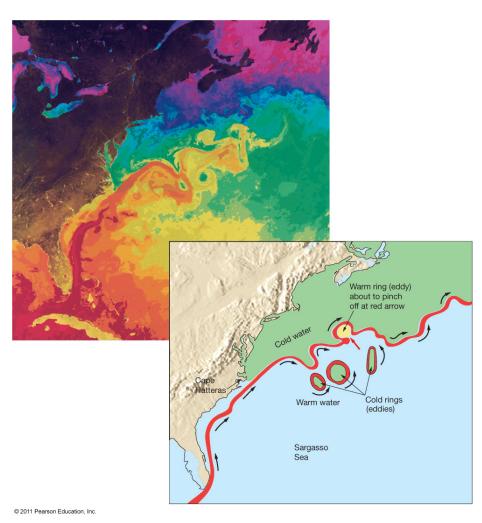




- as a western boundary current, the Gulf Stream is narrow, deep, and fast
- its western margin is abrupt, while its eastern margin is less defined
- it gradually merges eastward with the Sargasso Sea
- the Sargasso Sea is a body of water that circulates around the center of the North Atlantic Gyre
 - Sargassum is a floating marine algae that is particularly abundant in this area

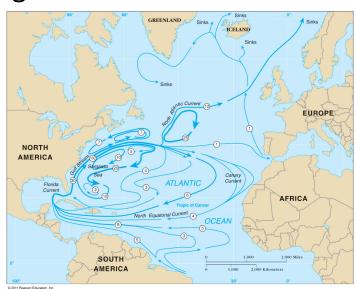
Rings of the Gulf Stream

- Meanders or loops may cause loss of water volume and generate:
 - Warm-core rings: warmer Sargasso Sea water trapped in loop surrounded by cool water. Narrow and shallow
 - Cold-core rings: cold water trapped in loop surrounded by warmer water. Wide and deep. Vertical mixing, with consequent upwelling and abundance of life
- Unique biological populations



Other North Atlantic Currents

- Labrador Current
 - comes from between Newfoundland and Greenland. Mixes with Gulf Stream to create fogbanks in the North Atlantic
- Irminger Current
 - west of Iceland
- Norwegian Current
 - west of Norway
- North Atlantic Current
 - crosses the Atlantic
- Canary Current
 - cold current that closes the gyre by rejoining the North Equatorial Current



Climate effects of North Atlantic Currents

North-moving currents – warm

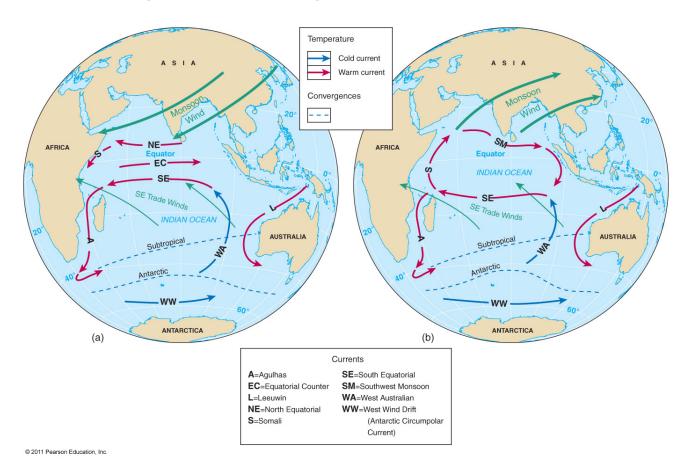
- the Gulf Stream warms the East coast of the United States and northern Europe
- North Atlantic and Norwegian Currents warm northwestern Europe

South-moving currents – cool

- Labrador Current cools eastern Canada
- Canary Current cools north African coast

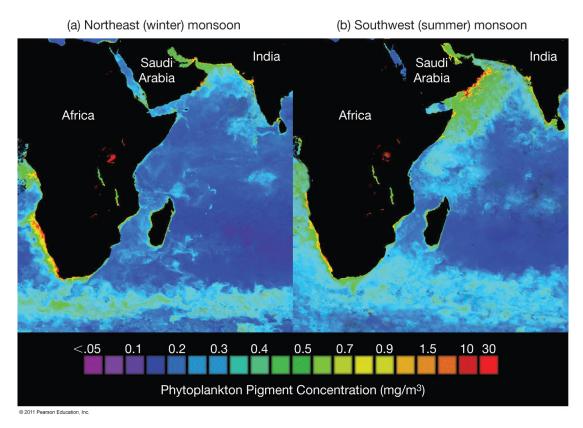
Indian Ocean Circulation

- Because of local geography, the Indian Ocean exists mostly in the southern hemisphere
- In winter, equatorial circulation is similar to Atlantic
- proximity to Himalayas and shape of coastline and ocean basin cause it to experience strong seasonal changes



- Seasonal reversals of winds over northern Indian Ocean
- Northeast monsoon (winter)
 - winds blowing from land: dry weather
- Southwest monsoon (summer)
 - winds blowing from ocean: heavy precipitation
- Affect seasonal land weather
- Affect seasonal Indian Ocean current circulation
- Affect phytoplankton productivity

Monsoons



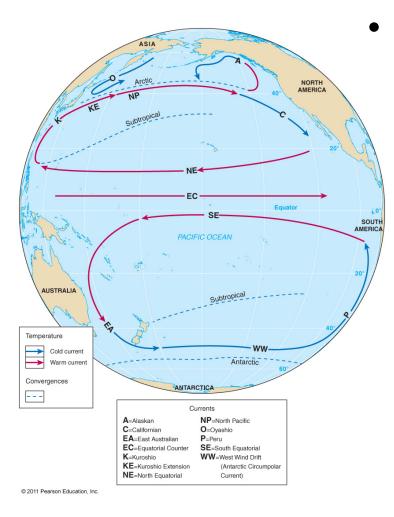
Indian Ocean Subtropical Gyre

Agulhas Current Australian Current

Leeuwin Current

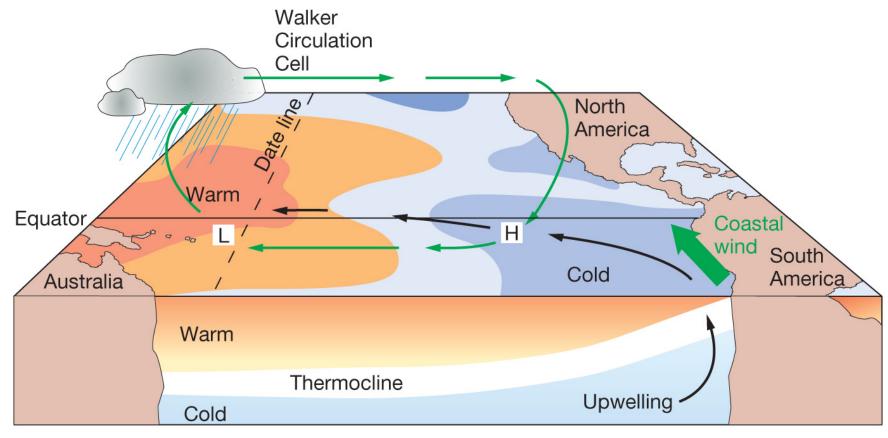
Pacific Ocean Circulation

- North Pacific Subtropical Gyre
 - Kuroshio
 - North PacificCurrent
 - CaliforniaCurrent
 - NorthEquatorialCurrent
 - AlaskanCurrent



- South Pacific Subtropical Gyre
 - East AustralianCurrent
 - AntarcticCircumpolarCurrent
 - Peru Current
 - SouthEquatorialCurrent
 - EquatorialCounterCurrent

Normal Conditions in the Pacific Ocean (Walker Circulation)



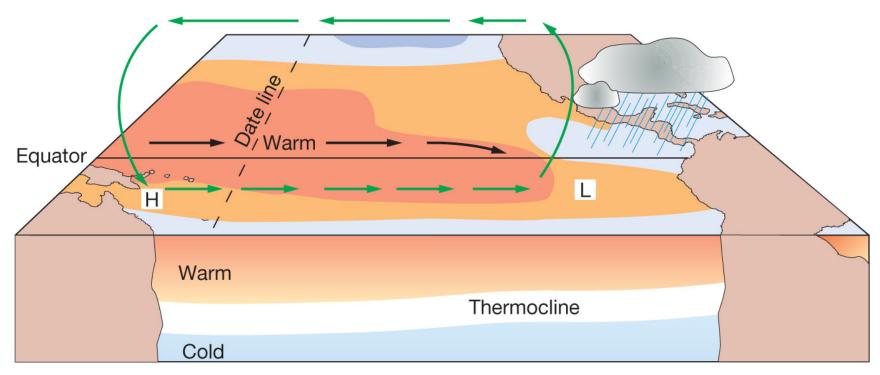
(a) Normal conditions

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Atmospheric-Ocean Connections in the Pacific Ocean

- Walker Circulation Cell normal conditions
 - Air pressure across equatorial Pacific is higher in eastern Pacific
 - Strong southeast trade winds
 - Pacific warm pool on western side of ocean
 - Thermocline deeper on western side
 - Upwelling off the coast of Peru

ENSO (El Niño – Southern Oscillation) Conditions in the Pacific Ocean



(b) El Niño conditions (strong)

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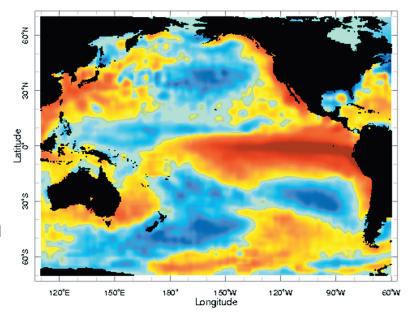
El Niño – Southern Oscillation (ENSO)

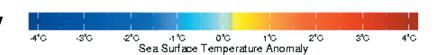
(a) Jan 1998

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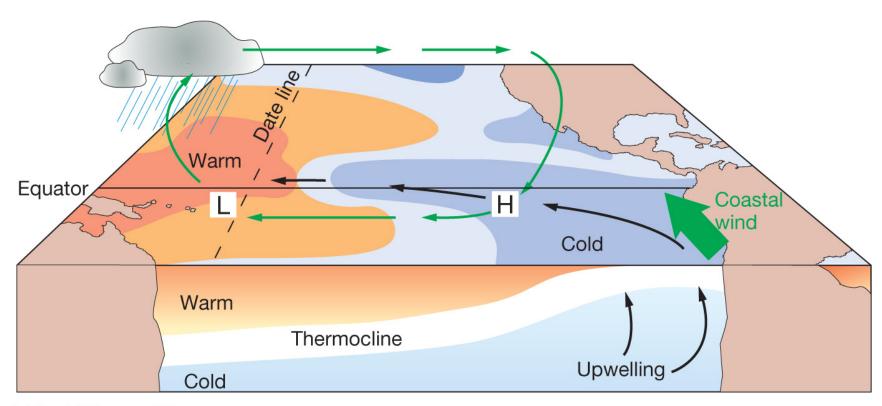
Walker Cell Circulation disrupted

- High pressure in eastern Pacific weakens
- Weaker trade winds
- Warm pool migrates eastward
- Thermocline deeper in eastern Pacific
- Downwelling
- Lower biological productivity
 - Peruvian fishing suffers





La Niña Conditions

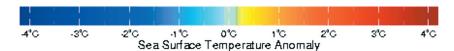


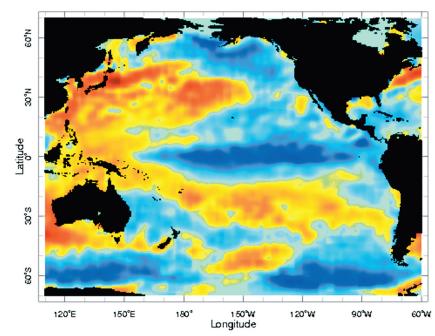
(c) La Niña conditions

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La Niña – ENSO Cool Phase

- Increased pressure difference across equatorial Pacific
- Stronger trade winds
- Stronger upwelling in eastern Pacific
- Shallower thermocline
- Cooler than normal seawater
- Higher biological productivity

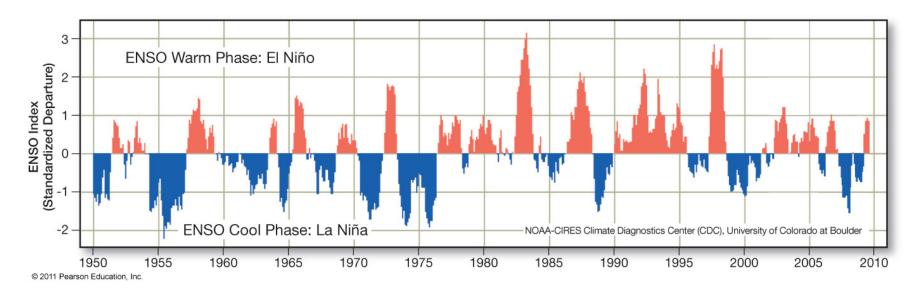




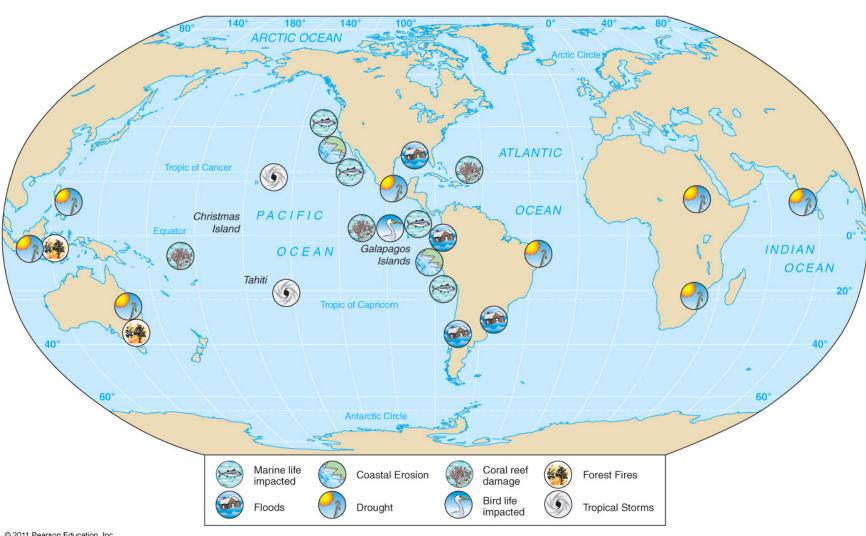
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How often do El Niño Events occur?

- El Niño warm phase about every 2–10 years
- Highly irregular
- Phases usually last 12–18 months
- 10,000-year sediment record of events
- ENSO may be part of Pacific Decadal Oscillation (PDO)
 - Long-term natural climate cycle
 - Lasts 20–30 years

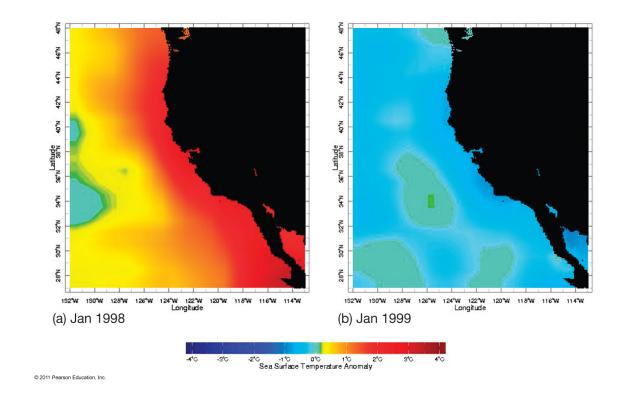


ENSO has Global Impacts



Notable ENSO Events

- 1982 1983
- 1997 1998
- Flooding, drought, erosion, fires, tropical storms, harmful effects on marine life
- Unpredictable



Predicting El Niño Events

- Tropical Ocean Global Atmosphere (TOGA) program
 - started in 1985
 - Monitors equatorial South Pacific
 - System of buoys
- Tropical Atmosphere and Ocean (TOA) project
 - Continues monitoring
- ENSO still not fully understood

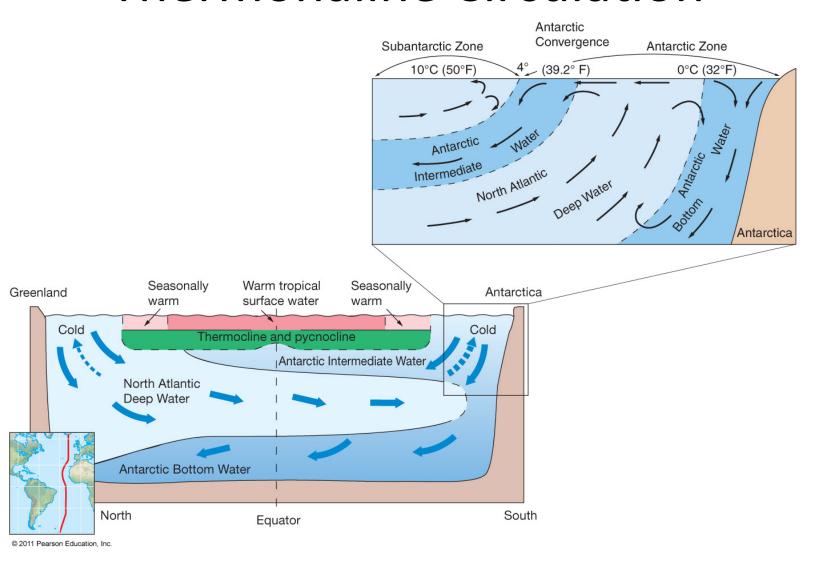
7.5 – How Do Deep-Ocean Currents Form?

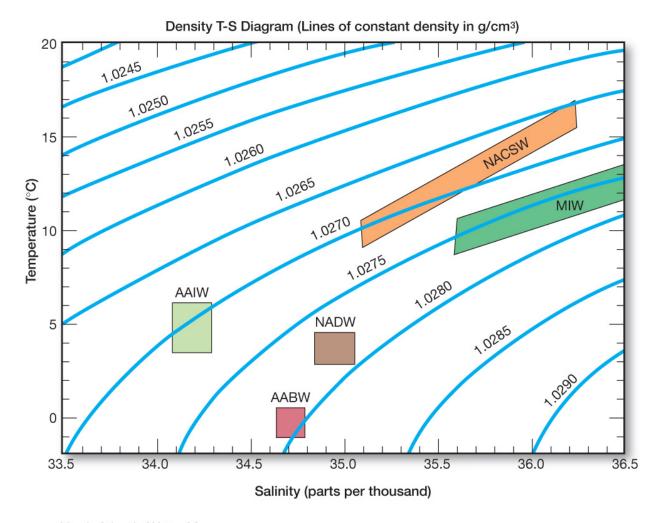
- Occur below the pycnocline
- As a consequence, they effect 90% of ocean waters
- They are created by variations in density of the ocean water
- This circulation system is then called thermohaline (temperature and salinity of ocean waters control their density)
- Thermohaline circulation is very slow: 10 to 20 km/yr

Origin of Thermohaline Circulation

- Originates in high latitude surface ocean
- Cooled, now dense surface water sinks and changes little
- Formation of sea ice increases the salinity, hence the density, of the water left behind, which starts to sink
- Deep-water masses can be identified on a temperature—salinity (T—S) diagram
 - Identifies deep water masses based on temperature, salinity, and resulting density

Thermohaline Circulation





North Atlantic Water Masses:

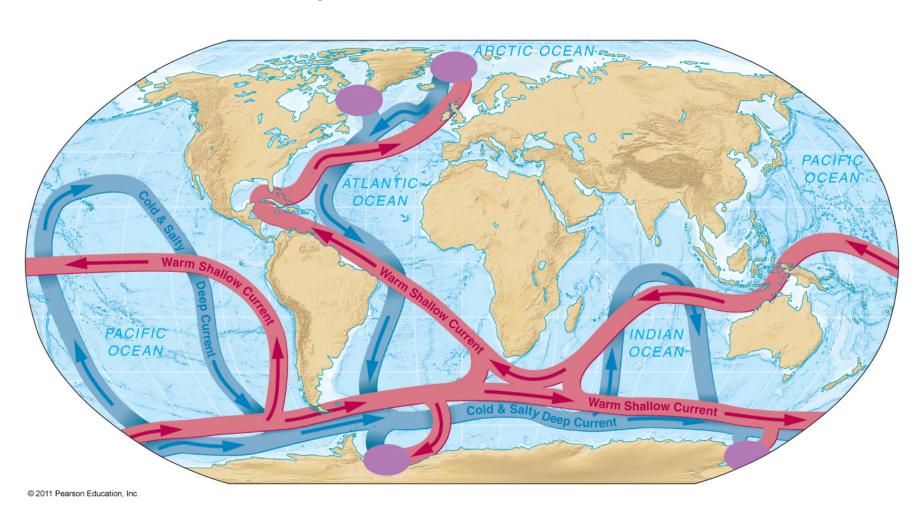


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Thermohaline Circulation

- Some deep-water masses
 - Antarctic Bottom Water
 - North Atlantic Deep Water
 - -Antarctic Intermediate Water
 - -Oceanic Common Water
- Cold surface seawater sinks at polar regions and moves equatorward

Conveyor Belt Circulation



- Dissolved oxygen in deep water
 - examples from the Cretaceous (greenhouse time)

- Conveyor-belt circulation and climate change
 - Plate tectonics changes geography
 - Climate changes affect deep-water circulation

7.6 – Can Power from Currents Be Harnessed as a Source of Energy?

- Currents carry more energy than winds
- Florida–Gulf
 Stream Current
 System
- Underwater turbines
 - Expensive
 - Difficult to maintain
 - Hazard to boating

